# Evaluation of Bank Storage Using Pressure and Solute Propagation





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As a requirement in full for the degree of Doctor of Philosophy in the School of the Environment, Flinders University of South Australia

## 2014



## **Table of Contents**

Table of Contentsi		
List of Figures	iv	
List of Tables	vii	
Summary	viii	
Declaration of (	Originalityx	
Co-authorship .	xi	
Acknowledgem	entsxii	
1 Introduction	on 1	
1.1 The r	esearch problem1	
1.2 Resea	arch aim3	
1.3 Struc	ture of this thesis	
2 Propagatio	on of solutes and pressure into aquifers following river stage rise	
2.1 Intro	duction	
2.2 Analy	rtical solution development	
2.3 Nume	erical simulations	
2.3.1	Model conceptualization and set up13	
2.3.2	Rapid increase in river stage17	
2.3.3	Sensitivity analysis for rapid increase in river stage	
2.3.4	Sensitivity to hydrological processes	
2.4 Exam	ple application to field data	
2.4.1	Field setting 29	

2.4	4.2 Determining travel times and travel distances	
2.5	Discussion	
2.6	Conclusions	35
3 Rela	elative rates of solute and pressure propagation into he	terogeneous alluvial aquifers
following	ng river flow events	
3.1	Introduction	
3.2	Methodology	41
3.2	2.1 Analytical solutions	
3.2.	2.2 Numerical simulations	43
3.3	Modelling Results	46
3.3	3.1 Process of pressure and solute propagation	46
3.3	3.2 Effects on travel time metrics	49
3.3	3.3 Comparison to analytical solutions	54
3.3	3.4 Estimating hydraulic conductivity and flux	56
3.3	3.5 Identifying heterogeneity from travel time ratio	os60
3.4	Field application	63
3.4	4.1 Site Description	63
3.4	4.2 Analysis of field data	64
3.5	Discussion	67
3.6	Conclusions	70
4 Infl	fluence of hydraulic gradient on bank storage exchange	e, penetration distance and
return time		
4.1	Introduction	73

4.2	Methods75
4.2.1	Analytical Solutions
4.2.2	Numerical Simulations77
4.2.3	Initial Conditions and Metrics 78
4.3	Results79
4.3.1	Process of Exchange
4.3.2	Aquifer and Wave Properties, Exchange and Return Time
4.3.3	River Shape and Aquifer Penetration 89
4.4	Discussion
4.5	Conclusions
5 Conc	lusions, Research Contribution and Future Work
Appendix	A Theoretical development of solute travel time 100
Appendix	B Analytical solutions for pressure propagation through a clogging layer 103
Appendix	C Synoptic sampling results from the Cockburn River 104
Appendix	D Time series data from the Cockburn and Mitchell Rivers 107
Appendix	E Groundwater chemistry and isotope results 109
Appendix	F Published conference proceedings 111
Reference	es

# List of Figures

Figure 2-1 Plot of <i>a</i> against <i>c</i> for commonly anticipated values in the context of bank
storage investigations (0 < c < 1)11
Figure 2-2 Model set up14
Figure 2-3 Comparison of analytical prediction and numerically simulated hydraulic heads
against time at $x = 10$ m and $z = 5$ m for the base case for a) the simulation period and b)
early time when $t_p$ occurs
Figure 2-4 Propagation of a) hydraulic head (m), b) Darcy flux (m d <sup>-1</sup> ) and c) solute
concentration into an aquifer due to a 1 m river stage rise18
Figure 2-5 Stage increase-induced variation in a) hydraulic head (m), b) Darcy flux (m $d^{-1}$ )
and c) solute concentration with respect to time at $x = 10$ m and $z = 1$ , 5 and 9 m
Figure 2-6 Sensitivity of solute and pressure travel time metrics to variation in $K$ , $b$ , $S_{y}$ , and
Н21
Figure 2-7 Sensitivity of solute travel time and travel distance to regional hydraulic
gradients at a) at x = 5 m and z = 5 m with K = 4.32 m d <sup>-1</sup> , b) at x = 10 m and z = 5 m with K =
4.32 m d <sup>-1</sup> , and c) at $x = 10$ m and $z = 5$ m with $K = 43.2$ m d <sup>-1</sup>
Figure 2-8 Percentage difference between the base case (zero gradient) time travel metrics
and travel time metrics for simulated positive (flow towards the river, gaining) and negative
(flow away from the river, losing) regional hydraulic gradients24
Figure 2-9 Hydraulic head and concentration plotted against time at $z = 5$ m, $x = 5$ and 10 m
for a 50 day stage increase and the base case (infinite stage increase)26
Figure 2-10 Percentage difference between the base case (full penetration, or $P = 1$ , river
bed at $z = 0$ m) travel time metrics and travel time metrics for partial river penetration P =
0.25 (river bed at $z = 7.5$ m) and P = 0.5 (river bed at $z = 5$ m) at observation points at $x = 10$
m and z = 2.5 m, 5 m, and 7.5 m27
Figure 2-11 Percentage difference between instantaneous rate of stage increase and
normalized pressure travel time for a representative range of aquifer parameters

Figure 2-12 Continuous pressure and EC measurement throughout a flood event in
September 2011 for the Cockburn River and two adjacent monitoring bores
Figure 3-1 Conceptual models of heterogeneity in alluvial aquifers, a) vertical clogging layer
and b) horizontal sand string
Figure 3-2 Comparison of numerically simulated a) head and b) solute propagation into a
heterogeneous aquifer with a 20 m thick clogging layer to two end-member homogenous
aquifers
Figure 3-3 Spatial distribution of solute and pressure propagation into aquifers containing 1
m thick sand strings (white lines) with increasing contrast in string and aquifer hydraulic
conductivites (numerical simulations)
Figure 3-4 Comparison of numerically simulated pressure ( $t_p$ ) and solute ( $t_s$ ) travel times
and their ratio $(t_s/t_p)$ at increasing distances from the river bank for homogenous aquifers
(solid lines) with low ( $K_1$ = 8.64 m d <sup>-1</sup> ) and high ( $K_2$ = 77.76 m d <sup>-1</sup> ) hydraulic conductivity and
heterogeneous aquifers (dashed lines) with either a)- c) clogging layer 20 m thick or d)-f)
sand string 1 m thick ( $H = 5 \text{ m}, K_2/K_1 = 9$ )
Figure 3-5 Comparison of numerical results for a) pressure travel times and b) solute travel
times as a function of hydraulic conductivity contrast for different sand string thicknesses
$(b_2)$ ( $K_2 = 77.76 \text{ m d}^{-1}$ , $H = 5 \text{ m}$ , $x = 30 \text{ m}$ )
Figure 3-6 Comparison of pressure and solute travel times obtained from analytical
solutions (equivalent homogenous, equations (3.1) and (3.2) and heterogeneous (equation
B1)) with travel times obtained from numerical simulations of heterogeneous aquifers 54
Figure 3-7 Comparison of a) hydraulic conductivities calculated with equation 3.1 ( $t_p$ ) or
equation 3.2 ( $t_s$ ) and b) fluxes from solute and pressure travel times obtained from
homogeneous aquifer simulations to actual (hydraulic conductivity) or model-derived (flux)
values
Figure 3-8 Ratios of a), b) estimated to average hydraulic conductivity and c), d) estimated
to simulated flux from river to aquifer for clogging layer and sand string scenarios

Figure 3-9 Numerically simulated solute and pressure travel times and their ratio at $x = 30$
m for a) clogging layers and b) sand strings of varying thickness across a range of hydraulic
conductivity ratios ( $K_2$ = 77.76 m d <sup>-1</sup> , $H$ = 5 m)61
Figure 3-10 Field data for May 2012 for piezometers at 15 m (GW1) and 30 m (GW2) from
the river bank64
Figure 4-1 Conceptual model of river – aquifer system75
Figure 4-2 Velocity vectors from the simulation of a wave 1 m high with a period of 5 days
under a hydraulic gradient of 0.01 depicting the process of river water movement into and
out of an aquifer80
Figure 4-3 Total volume of water exchange per unit width of aquifer and return times
obtained from analytical solutions for a range of hydraulic conductivity (K), wave height (H)
and wave period values at a range of hydraulic gradients83
Figure 4-4 Dependency of total exchange per unit width of aquifer, penetration distance,
and 90% return time on variation in hydraulic gradient and hydraulic conductivity (K), wave
height (H) and period, and dispersivity (A)85
Figure 4-5 Percentage error in analytical results compared to numerical results for a) total
flux and b) 90% return time86
Figure 4-6 10%, 50% and 90% return times after commencement of river stage rise for
variation in hydraulic conductivity (K), wave height (H) and period (d), dispersivity (A), and
partial penetration with varying aspect ratios (AR)
Figure 4-7 Vertical distribution of exchange along the wetted perimeter on entry to the
aquifer for a range of hydraulic conductivities at a) hydraulic gradient = 0 and b) hydraulic
gradient = 0.01
Figure 4-8 Dependency of a) total exchange per unit width of aquifer, b) horizontal
penetration distance, c) vertical penetration distance, and d) return time on variation in
hydraulic gradient and partial penetration90

Figure 4-9 Close up of the spatial variation in penetration distance of the 0.5 concentration	า
contour for a fully penetrating river (FP, black line), aspect ratio (AR) of 3 (red line) and 9	
(blue line) with no hydraulic gradient present	€
Figure 4-10 Comparison of river flow events in a) 2012 and b) 2013 and the groundwater	
level (solid red line) and specific electrical conductivity (EC, dashed red line) at a field site	
adjacent to the Mitchell River in tropical north Queensland, Australia	<del>)</del> 4

## **List of Tables**

Table 2-1 Model parameters for base case simulation.	16
Table 3-1 Parameters used for generic model simulations.	45
Table 4-1 Parameter variation.	78

#### **Summary**

Bank storage is the process of river water mixing with near-river groundwater as a result of an increase in river stage due to a flow event. Such mixing causes temporal and spatial variation in near – river groundwater chemistry. However, the extent of the interaction is poorly defined. The extent of the interaction has important ramifications for biogeochemical cycling, contaminant mixing and degradation, and resource assessment techniques that differentiate between surface water and groundwater reservoirs. Previous assessments of bank storage have primarily relied on hydraulic methods, particularly pressure propagation, and chemistry measurements with limited temporal resolution. This work aimed to evaluate the relative rates of solute and pressure propagation and develop new assessment techniques for bank storage in a variety of hydrogeological environments.

In contrast to pressure propagation into homogeneous aquifers in response to river stage rise, the relationships between water propagation and aquifer properties were not well understood prior to this study. Practically, water movement is most readily measured using a conservative solute or tracer. Numerical assessment of a new analytical relationship between solute and pressure travel times and distances and aquifer and flow event characteristics determined that the solution may be used in variably saturated aquifers with errors generally less than 30%. In homogeneous aquifers the ratio of solute to pressure travel time is independent of hydraulic conductivity. Consequently, under certain hydrological conditions time series measurement of pressure and a solute (or proxy) and computation of pressure and solute travel times enables a first-order estimate of aquifer properties and the lateral extent of river water penetration into an aquifer.

In homogeneous systems river stage rise causes pressure to propagate faster and further into an aquifer than water (or solutes). Numerical testing of two conceptual models of alluvial heterogeneity indicated that pressure and solute propagation are unequally affected by aquifer heterogeneity. Hence, under certain conditions, substantial solute change can be recorded in an aquifer before substantial pressure change. This may be identified by computing a solute travel time less than a pressure travel time. Flux estimates obtained from solute travel times using homogeneous solutions were determined to be more accurate than estimates obtained from pressure data. The error in estimates derived from pressure data was proportional to the contrast in hydraulic conductivity in a system.

Theoretical investigations of bank storage have not systematically quantified the influence of the hydraulic gradient between aquifer and river. In this work analytical and numerical techniques demonstrated that variation in the hydraulic gradient influences bank storage exchange, penetration distance and residence time, at a scale similar to substantial variation in hydraulic conductivity, wave height and period, dispersivity, and river partial penetration. Consideration of the hydraulic gradient is therefore integral to quantitative assessments of exchange.

Simultaneous measurement of pressure and solutes at high temporal resolution within rivers and adjacent aquifers is a useful technique for improving understanding of the spatial and temporal extent of river – aquifer exchange during flow events. The utility of the theory relies on contrasting river and aquifer chemistries. Future work should consider the use of alternative tracers to test residence time distribution theories, and geostatistics, spatial imaging, and uncertainty techniques to further understand the influence of heterogeneity.

## **Declaration of Originality**

I certify that this thesis does not incorporate, without acknowledgment, any material previously submitted for a degree or diploma in any other university; and that to the best of my knowledge and belief it does not contain any material previously published or written by another person except where due reference is made in the text.

Chani Welch

## **Co-authorship**

Chani Welch is the primary author on this thesis and all the enclosed documents. Chapters 2 to 4 were written as independent manuscripts in which the co-authors provided intellectual supervision and editorial comment. Neville I. Robinson completed the analytical derivation presented in Appendix A.

### Acknowledgements

First and foremost I would like to acknowledge the contributions of my two supervisors to this PhD. I am grateful to Peter Cook for his generosity with his many ideas and time, for sharing his passion for hydrogeological process understanding, and for expecting the best. I appreciate the numerous lengthy conversations with Glenn Harrington about hydrogeology, modelling, and various completely unrelated topics, which, as much as anything, helped me work things out for myself. I would also like to express my gratitude to Marc Leblanc for coordinating all aspects of field work in the Mitchell River and being continually passionate about tropical hydrology, and Roki for the fresh perspective he brought at just the right time.

I gratefully appreciate assistance with field work provided by Jordi Batlle Aguilar, Roger Cranswick, Saskia Noorduijn, Rolf Kipfer, Michelle Irvine, Cameron Wood, Nick White, Lawrence Burke, Mark Trigg, Vincent Post, Nick Rockett, Tony Forsyth, Joel Bailey, and the NSW Office of Water.

Financial assistance consisted of an Australian Postgraduate Award and a scholarship from the National Centre for Groundwater Research and Training (NCGRT). Additional funding for this research was provided by the NCGRT, an Australian Government initiative, supported by the Australian Research Council and the National Water Commission. Funding for the installation of piezometer transects adjacent to the Cockburn River was provided through the SuperScience program. The Australian Wildlife Conservancy – Brooklyn provided site access and general assistance during field work on the Mitchell River.

On a personal note, I would like to thank my Adelaide family for the companionship, camping, food, and wine that have made the past few years so enjoyable, and my family and friends who stuck around for the journey, helping me maintain perspective.

Chani Welch 7 February, 2014

### **1** Introduction

#### **1.1** The research problem

Quantifying the exchange of water and solutes between rivers and aquifers has become crucial in many water-limited environments as climatic conditions and human consumption stress surface water and groundwater systems simultaneously (Baillie et al., 2007). This situation has highlighted the interconnection between the two resources, and historical over-allocation of water in some areas (Nevill, 2009; Sophocleous, 2002). A variety of methods are available to quantify water fluxes between aquifers and rivers, but many are limited by the localised nature of the measurements, and spatial heterogeneity in hydrogeological properties (Kalbus et al., 2006). This can result in highly variable and uncertain volume estimates, and, where such values are subsequently used for management purposes, potential sub-optimal allocation of resources. In addition, estimates made using hydraulic and chemistry based methods often produce conflicting results (Kirchner, 2003).

Characterising the river-aquifer exchange processes occurring in a system is the first step to appropriately quantifying groundwater discharge or recharge. Exchange processes occur over a continuum of timescales and include hyporheic exchange (minutes to weeks), parafluvial flow (hours to months) and bank storage (hours to years). Hydrological, geological, and geomorphological controls on the significance of these processes vary throughout time and space, as do the drivers for the exchange (Winter, 1998). These short to medium term mixing processes occur, and therefore need to be considered, in the context of regional processes that occur over decades to millennia as a result of aquiferwide recharge/discharge dynamics.

Chemical mass balance flux quantification methods are considered most appropriate on a scale potentially useful to water resource managers due to their ability to integrate processes over larger scales (Kalbus et al., 2006). However, interpretation of chemistry data

using simple models requires that processes are lumped, and hence, results can be highly dependent on the conceptual model and parameterisation. Lack of consideration of hyporheic exchange, for example, can significantly affect estimates of groundwater discharge to rivers (Cook et al., 2006). Furthermore, river aquifer exchange processes mix river water with groundwater. In many cases this creates a temporally and spatially variable zone of water around a river with a chemistry that is distinct from the wider aquifer. Lack of consideration of this temporal variability has rarely been considered but can lead to significant errors in estimates of groundwater discharge to rivers (McCallum et al., 2010). Techniques that exploit temporal chemistry changes to determine, for example, travel times from rivers to aquifers, are increasing as measurement technologies improve, but existing analysis methods such as deconvolution (e.g., Cirpka et al., 2007; Vogt et al., 2010) and principal component analysis (Lewandowski et al., 2009; Page et al., 2012) cannot explicitly identify the influences of individual aquifer properties or processes.

Temporal changes in groundwater chemistry have many drivers. In aquifers connected to rivers, flow events are considered to be one of the main drivers. As the river level rises above that of the adjacent groundwater, river water moves into the aquifer, and mixes with the existing groundwater. As the flood wave passes, this mixture of water returns to the river. This process is termed bank storage (Todd, 1956). Bank storage occurs regardless of whether a river is gaining or losing water from the adjacent aquifer, as long as the river and aquifer are hydraulically connected.

Analytical solutions that relate pressure propagation as a function of river stage rise to aquifer properties, and from that estimate flux, have long been available for homogeneous systems (Cooper and Rorabaugh, 1963) and for rivers with clogging layers (Hall and Moench, 1972; Hantush, 1965). However, pressure change alone cannot be directly correlated to the extent of water movement into an aquifer. Conventional theory indicates that pressure propagates further and faster than water or any solutes it contains. However,

2

explicit relationships between solute propagation, river stage rise, and aquifer properties have not been described in analytical solutions. Consequently, solute data has been underutilised. Similarly, numerical investigation of the process of bank storage has been predominantly hydraulic, and in homogenous systems, as pressure data is more readily available for model calibration and simulating solute transport is more numerically intensive. Field assessment of bank storage has also predominantly used hydraulic methods, or measurements of river chemistry obtained at low temporal resolution. Systematic assessment of the movement of water and solutes during bank storage was required to facilitate exploitation of increasingly available high temporal resolution solute data for the purposes of improved management of surface water and groundwater resources.

#### **1.2 Research aim**

The aim of this research was to increase bank storage process understanding in general, evaluate the relative rates of pressure and solute propagation during bank storage, and develop new techniques for assessment of river – aquifer exchange during flow events. The research was based on two hypotheses:

- River stage rise and fall induces a predictable variation in near-river groundwater chemistry, the spatial and temporal extent of which is dependent on key aquifer properties and the degree of stage change; and
- ii. Continuous measurement of a solute (or proxy) in addition to pressure in near-river groundwater will assist with the determination of additional aquifer properties (compared to solely measuring pressure) and the conceptual model of the riveraquifer interface.

In order to address these hypotheses this work aimed to:

use analytical and numerical methods to explore relationships between water
(which can be represented by a generic conservative solute) and pressure travel

time and distance and aquifer and wave characteristics in variably saturated aquifers with a wide range of characteristics and conceptual models;

- provide a preliminary assessment of the influence of heterogeneity structures on the relative rates of solute and pressure propagation identified for homogenous aquifers using transient numerical flow and transport simulations;
- examine the relationships of bank storage exchange, penetration distance and return time to aquifer properties, river – aquifer conceptual models, and hydraulic gradients; and
- verify the practical application of theoretical findings using data collected in distinct hydrogeological environments. Field sites were instrumented in semi-arid northern New South Wales on the Cockburn River and tropical north Queensland on the Mitchell River.

#### 1.3 Structure of this thesis

This thesis consists of a broad overview (Chapter 1), three pieces of work published in or submitted to international peer-reviewed journals (Chapters 2 - 4) and overarching conclusions of the research, including the research contribution and recommendations for further work (Chapter 5). The three manuscripts included are:

- Welch, C., P. G. Cook, G. A. Harrington, and N. I. Robinson (2013), Propagation of solutes and pressure into aquifers following river stage rise, *Water Resources Research*, 49, 5246–5259, doi:10.1002/wrcr.20408 [Chapter 2];
- (2) Welch, C., G. A. Harrington, M. Leblanc, J. Batlle-Aguilar, and P. G. Cook (2014), Relative rates of solute and pressure propagation into heterogeneous alluvial aquifers following river flow events, *Journal of Hydrology*, 511, 891-903, doi: 10.1016/j.jhydrol.2014.02.032 [Chapter 3]; and

(3) Welch, C., G. A. Harrington, and P. G. Cook (under review), Influence of hydraulic gradient on bank storage exchange, penetration distance and return time, submitted to *Groundwater* [Chapter 4].

Supplementary information for Chapters 2-4 is contained in appendices, as are conference papers which resulted directly from this research.